

Supplementary Information section for:

A simple method for evaluating photochemical dependencies in the atmosphere

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S1 Box model simulations of photo-chemical budgets for OH, RO_x and O_x

Computed diurnal evolutions of key species are shown in Figure 1 of the main text. Box model simulations of gas-phase chemical species in the conditions of remote atmosphere were performed in order to gain an understanding of behavior of species studied in this work. The model-calculated budgets for HO, RO_x (OH + HO₂ + RO₂) and O_x (O₃ + NO₂) are shown in Figure 1. The budgets are shown for the fifth day of the simulation so as to give other non-fixed intermediates enough time to reach near steady state conditions. The loss of OH averaged over the fifth diurnal cycle is dominated by CO (75%) with a minor contribution from CH₄ (13%). Less than 5% of the OH loss is due to NMHCs and oxygenated organic intermediates in this pristine environment. In most cases, these OH losses produce peroxy radicals, HO₂ and RO₂. The OH production terms include HO₂+NO (35%), HO₂+O₃ (34%), and O¹D+H₂O (26%), with only a minor contribution from hydrogen peroxide (H₂O₂) photolysis (3.7%). The loss of RO_x averaged over the fifth diurnal cycle is dominated by HO₂ self-reaction (64%) with a minor contribution from CH₃O₂+HO₂ (26%). The production of RO_x is dominated by O¹D+H₂O (57%) with contributions from HCHO photolysis (20%) and H₂O₂ photolysis (8.1%). The production of O_x over the fifth diurnal cycle is dominated by HO₂+NO (73%) and RO₂+NO (24%). Loss of O_x is dominated by HO₂+O₃ (63%) with contributions from O¹D+H₂O (24%) and OH+O₃ (12%). The net rate of change over the fifth simulation day is small, with near balance between production and loss (-0.12 ppbv/day).

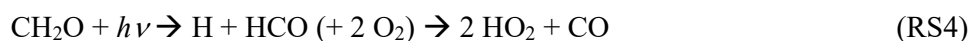
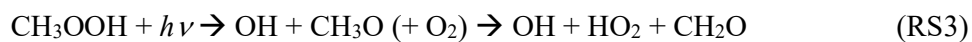
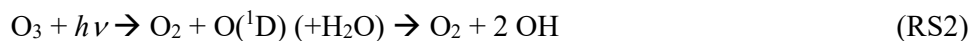
Direct evaluation of the modeled HO_x, RO_x, and O_x budgets with observations is difficult because production and loss rates are not measured directly. We note that the small net production of O_x is consistent with the observation that O₃ concentrations were relatively steady during TOPSE. However, we also note that the calculated budget of HO₂ radicals is problematic *vis-à-vis* H₂O₂ measurements (see Table 2), which are well below 1 ppb with average ~ 320 ppt and medians considerably lower, near 100 ppt. The model (see Fig. 1b)

suggests stronger H₂O₂ production, approx. 2.5 ppb/day at noon or, taken over a diurnal cycle, about 1 ppb per day. With this large production, the low measured concentrations point to an H₂O₂ lifetime of approx. 1/3rd of a day, much shorter than the diurnally averaged lifetimes estimated for H₂O₂ loss by photolysis (approx. 2 days) and by OH (approx. 13 days). This may be pointing to the importance of heterogeneous losses of H₂O₂ being dominant under these selected TOPSE conditions. The fact that measured CH₃OOH levels are comparable to H₂O₂ levels, while the gas-phase model predicts much smaller CH₃OOH/H₂O₂ ratios, is consistent with this explanation, given the higher solubility of H₂O₂. An alternative interpretation is that the box model overestimates HO₂ radical concentrations, the predicted H₂O₂ production of course being proportional to [HO₂]².

S2 Box model calculations of sensitivity coefficients used for LSFs

The model-calculated sensitivity coefficients are given in Table 3, can be negative, and have small ranges (shown as errors in Table 3) over the large variations explored. The detailed sensitivities can be understood easily, at least qualitatively, in terms of the known photochemistry occurring in the pristine troposphere, as detailed below.

J values: Photolysis reactions are the sources of most HO_x radicals. Based on Fig. 1c, the most important of these are:



Two HO_x radicals are produced in each case. This production is balanced (see Fig. 1c) by peroxy radical reactions (RS5-6) and to a lesser extent the OH termination reaction (RS5):





Neglecting the organic terms, the HO_x rate equation is

$$\frac{d[\text{HO}_x]}{dt} = 2J_1[\text{H}_2\text{O}_2] + 2J_2[\text{O}_3]f[\text{H}_2\text{O}] - 2k_5[\text{HO}_2]^2 - k_7[\text{OH}][\text{NO}_2] \quad (\text{Eq. S1})$$

where f is related to the probability of O(¹D) reacting with water rather than being quenched.

In low NO_x conditions, the last term is relatively small, so at steady state:

$$[\text{HO}_2] = \sqrt{\frac{J_1[\text{H}_2\text{O}_2] + J_2[\text{O}_3]f[\text{H}_2\text{O}]}{k_5}} \quad (\text{Eq. S2})$$

This square root dependence is well borne-out for peroxy radicals in Table 3. However, the stronger, nearly linear increase in OH is due to the increase in another photolysis reaction,

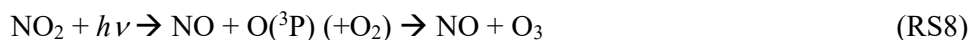


Table 3 shows that increasing J values result in an almost linear increase in the NO/NO₂ ratio, with the higher NO levels (at constant NO_x) then leading to OH reformation via the reaction:



Water (H₂O): Increased water concentrations lead to a larger fraction of O(¹D) reacting to form OH:



This reaction is a significant but not unique source of HO_x radicals, with photolysis of peroxides being at times comparable or even more important. Thus, the effect on HO₂ and RO₂ (the major HO_x constituents) falls below the square root dependence found, e.g., for J values, as is also clear from Eq. S2. The small negative effect on the NO/NO₂ ratio results from the reaction of NO with increased peroxy radical amounts.

Nitrogen oxides (NO_x): Increases in NO_x are seen to have little effect on the NO/NO₂ partitioning, or, conversely, model-predicted NO/NO₂ ratios are insensitive to the exact NO_x level. The availability of more NO_x does lead to more OH production, e.g., if added as NO:



but this is also a loss for HO₂, which is closely coupled to OH by the CO reaction:



To the extent that these two reactions contribute to HO_x partitioning, the HO₂/OH ratio should scale as CO/NO. Thus, higher NO_x levels lead to higher OH, to constant HO₂ values, and to negative RO₂ changes. This last effect results from the reaction:



being more important for RO₂ than for HO₂, due to the absence of the organic analogue to the reaction:



Adding NO_x as NO₂, provides some additional reactivity through ozone formation.

Ozone (O₃): Ozone affects radical production via direct photolysis (as already discussed for H₂O and *J* sensitivities), and affects the NO_x photostationary state:

$$\frac{[\text{NO}]}{[\text{NO}_2]} = \frac{J_5}{k_4[\text{HO}_2] + k_5[\text{RO}_2] + k_6[\text{O}_3]} \quad (\text{Eq. S3})$$

Of the three terms in denominator, HO₂ and RO₂ are negligible compared to O₃ for these conditions (see model simulations in Fig. 1). Thus the NO/NO₂ ratio is nearly proportional to *J*₅/O₃.

Carbon monoxide (CO): Increases in CO are seen to decrease OH substantially, in accord with Fig. 1 that shows their direct reaction to be a major part of the total OH budget. The decrease in OH results in lower oxidation rates for hydrocarbons, and thus fewer RO₂'s produced. The increase in the number of HO₂ is equal to the number (not the percent) decrease of RO₂ produced, at least within the rough approximations that OH losses are always accompanied by production of either HO₂ or RO₂, and these having similar lifetimes. Therefore, the total, HO₂+RO₂, remains constant, as does the NO/NO₂ ratio. Little if any additional ozone production or destruction would be expected by addition of CO.

Temperature (*T*): The rate constants for the important reactions $\text{OH} + \text{CH}_4$ and for $\text{NO} + \text{O}_3$ both increase strongly with temperature, resulting in faster RO_2 production and smaller NO/NO_2 ratios.

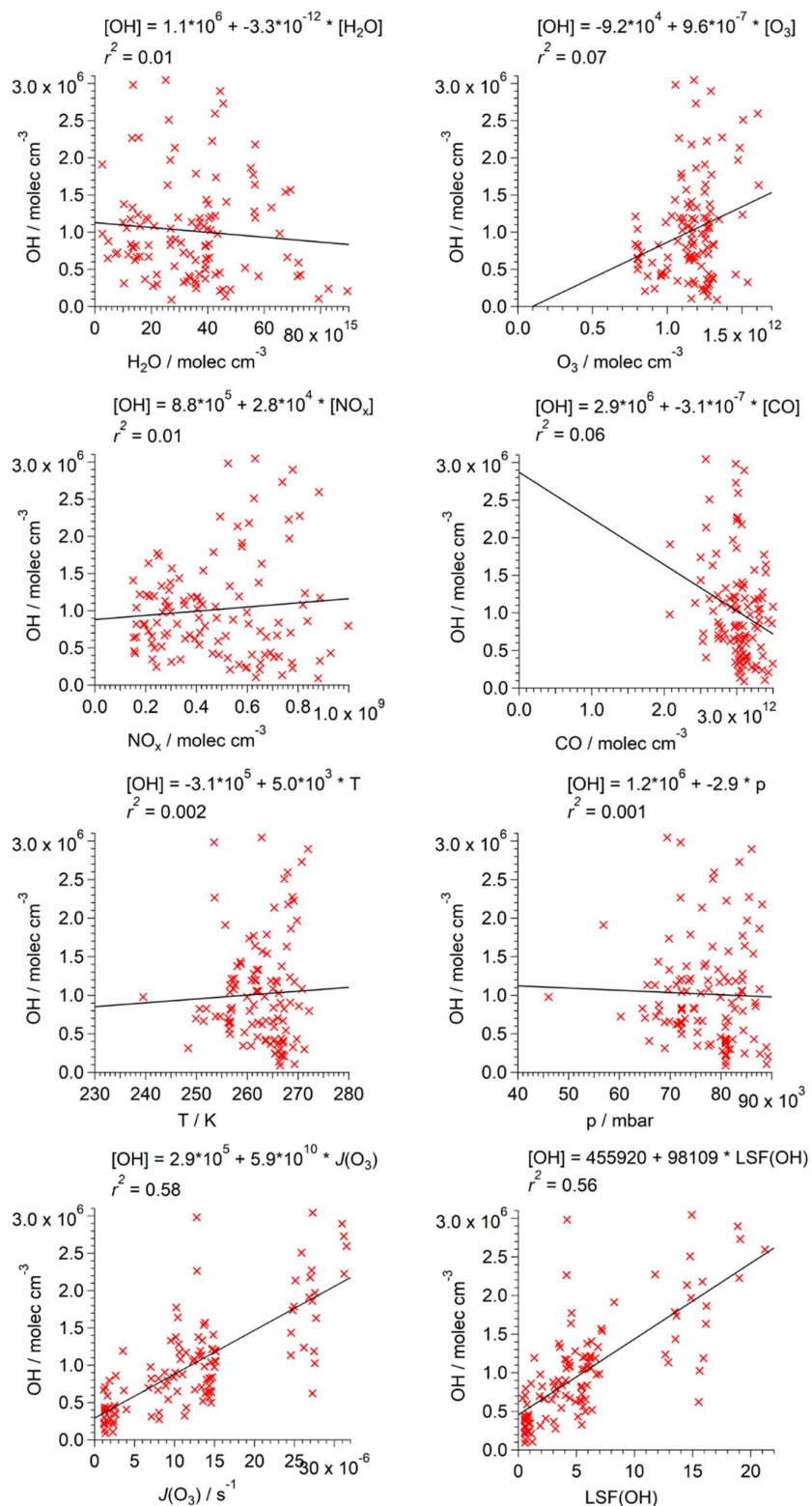


Figure S1. Correlations of OH with TOPSE-measured species (H₂O, O₃, NO_x and CO) and meteorological parameters (T and p) used in LSF. Also given are correlations of OH with J value ($\text{O}_3 + h\nu \rightarrow \text{O}_2 + \text{O}(^1\text{D})$) and LSF (bottom two panels) shown in Figure 3 of the main text, for the ease of direct comparisons.

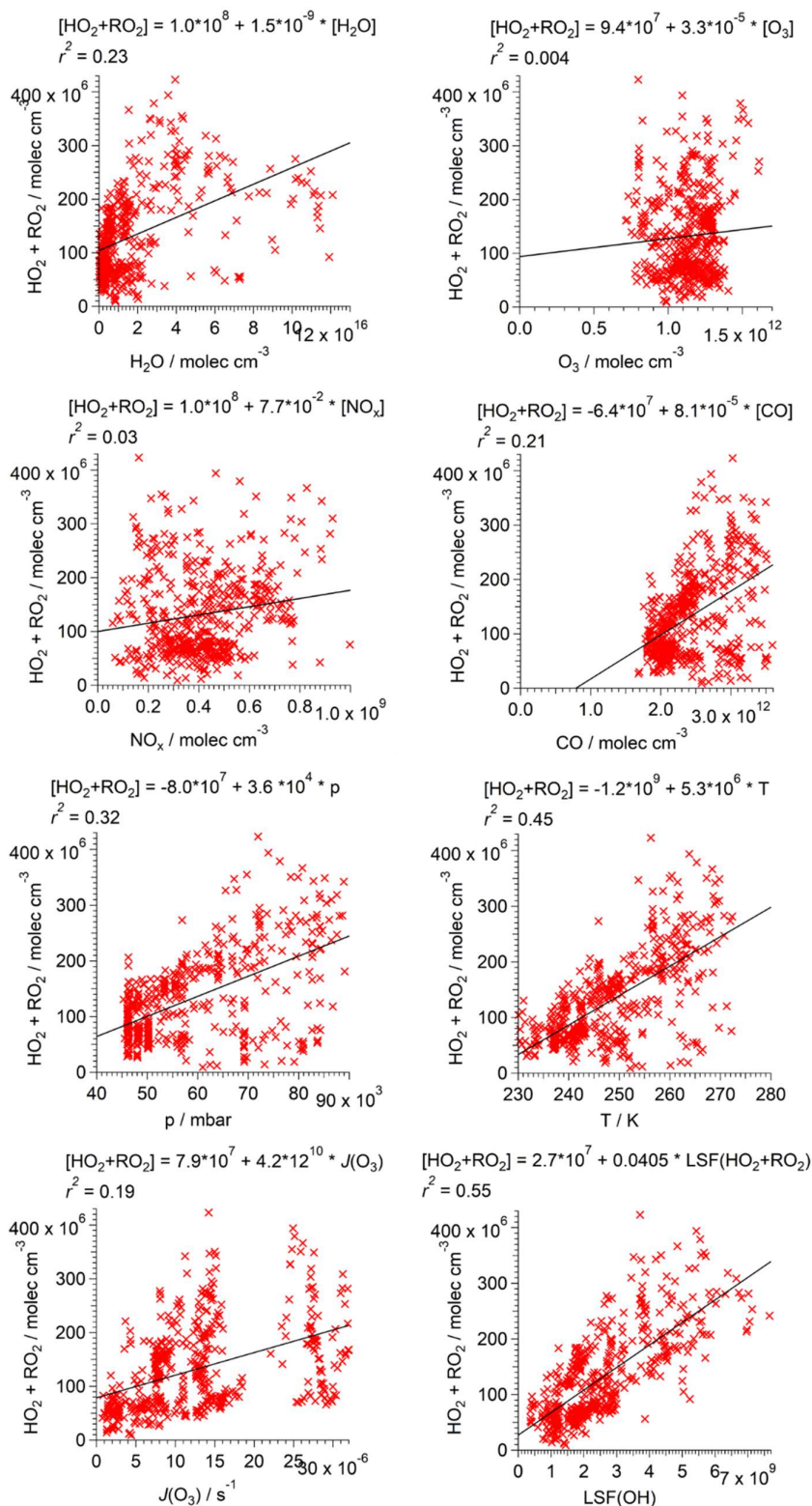


Figure S2. Correlations of $\text{HO}_2 + \text{RO}_2$ with TOPSE-measured species (H_2O , O_3 , NO_x and CO) and meteorological parameters (T and p) used in LSF. Also given are correlations of $\text{HO}_2 + \text{RO}_2$ with J value ($\text{O}_3 + h\nu \rightarrow \text{O}_2 + \text{O}(^1\text{D})$) and LSF (bottom two panels) shown in Figure 3 of the main text, for the ease of direct comparisons.

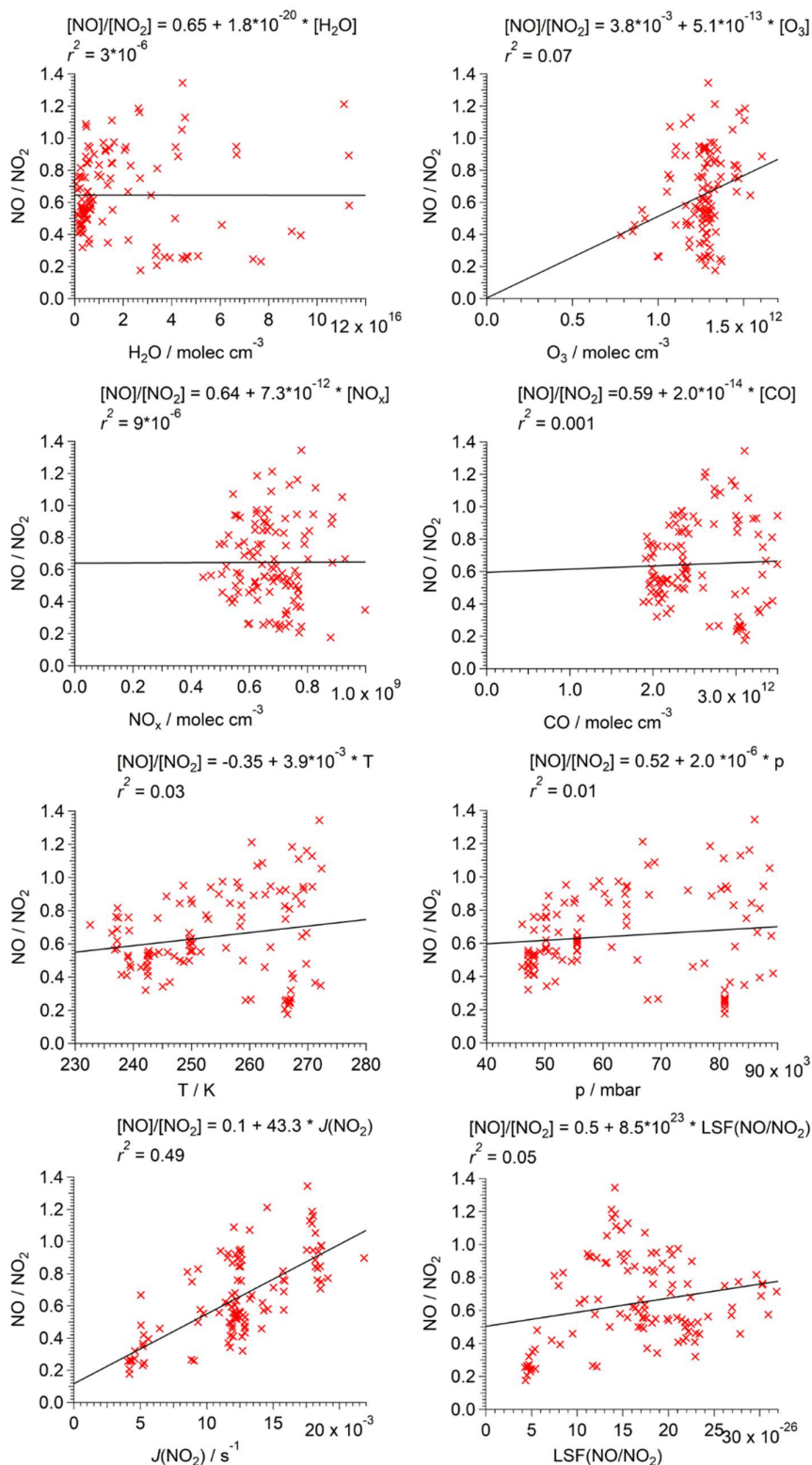


Figure S3. Correlations of NO/NO₂ ratio with TOPSE-measured species (H₂O, O₃, NO_x and CO) and meteorological parameters (T and p) used in LSF. Also given are correlations of NO/NO₂ with J value ($\text{NO}_2 + h\nu \rightarrow \text{NO} + \text{O}(^3\text{P})$) and LSF (bottom two panels) shown in Figure 3 of the main text, for the ease of direct comparisons.